

Partitioning of iron between perovskite, post-perovskite, and ferropericlase up to 154 GPa and 2000 K

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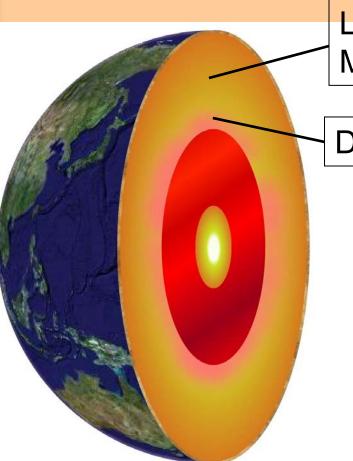
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SPring-8

Introduction



Lower (Mg,Fe)SiO₃ Perovskite (Pv) Mantle (Mg,Fe)O Ferropericlase(Fp)

D" layer Fp

Partition coefficient (K) of Fe-Mg between Pv/PPv and Fp

$$K = \left(\frac{X_{Fe}}{X_{Mg}}\right)^{Pv \text{ or } PPv} / \left(\frac{X_{Fe}}{X_{Mg}}\right)^{Fp}$$

 $X_{Fe} = FeO/(FeO + MgO)$ in molar ratio

Fe-Mg partitioning between these phases is important for understanding the chemical and physical properties (e.g. density, elastic velocity) of Earth's lower mantle.

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Partition coefficient (K) of Fe-Mg between Pv and Fp

Pressure dependency of K between Pv and Fp

K increases with pressure

Mao et al., 1997 Andrault, 2001; K>1 above 70-80GPa Determined by X-ray (unit cell volume)

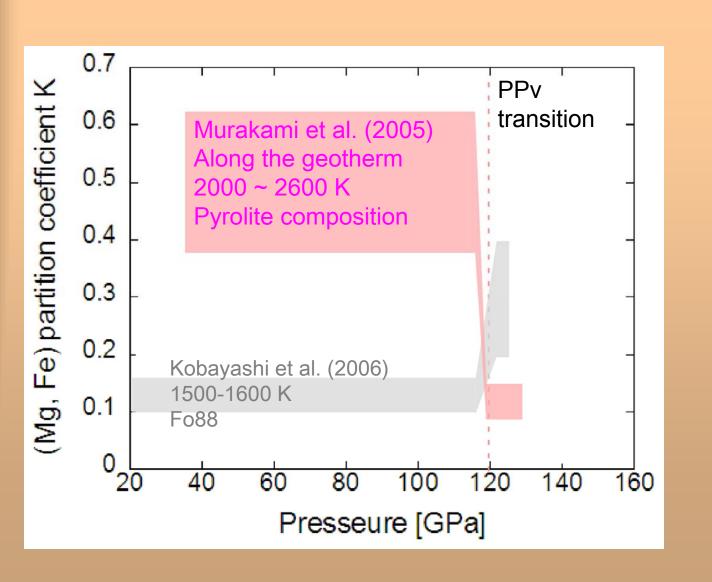
K is almost constant with pressure

Guyot et al., 1988; constant above 40GPa Kesson et al., 1991, 1998, 2002

Determined by ATEM



Partition coefficient (K) of Fe-Mg between Pv/PPv and Fp



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HS-LS transition in Magnesiowüstite

Iron may strongly prefer the LS-Mw (Badro et al., 2004)

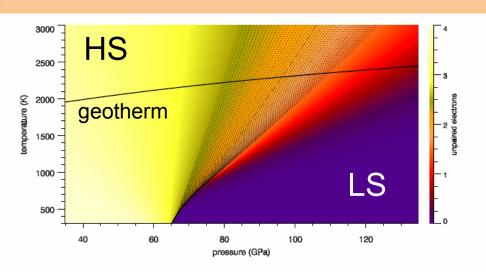
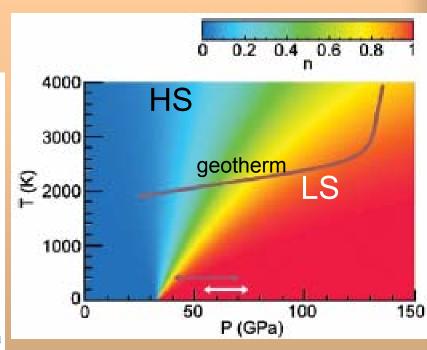


Figure 3. The average number of unpaired electrons for Mg_{0.83}Fe_{0.17}O in the pressure-temperature sector. The dotted contours show the region between 1.5 and 2.5 unpaired electrons. This region clearly widens with increasing temperature. In particular, the iron spin state varies smoothly along a lower mantle geotherm [Brown and Shankland, 1981] given by the solid line. Parameters used for this calculation are given in the text.



Sturhahn et al., 2005 GRL

Tsuchiya et al., 2006 PRL



Experimental procedure

Starting Material: San Carlos olivine (Mg_{0.88}, Fe_{0.12})₂SiO₄ Al-free system

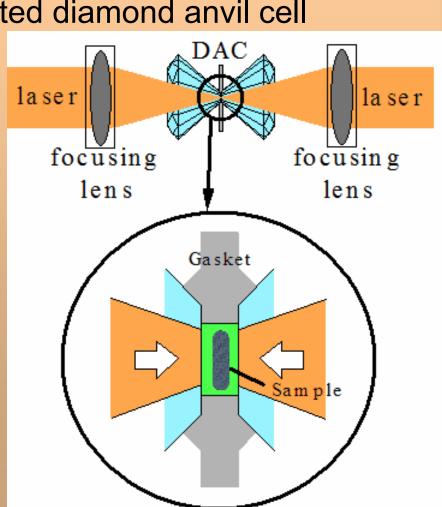
High P-T generation: Laser-Heated diamond anvil cell

Laser spot size: ~50 µm

Pressure medium: with or without NaCl

P measurement:
Ruby fluorescence
T2g Raman mode of Diamond
(Akahama et al., 2004)

T measurement: Spectroradiometric method (e.g. Boehler 2000)



Analytical methods

 Lattice volume measurement at ambient condition (Determined by X-ray)

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Pv : V_0^{Pv} = 162.51 + 6.00X (Å)

Mw : V_0^{Mw} = 74.78 + 6.65X (Å)

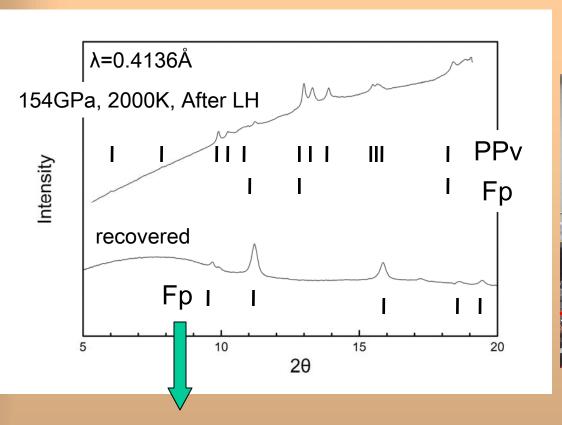
(X = FeO / (FeO + MgO) in molar ratio)
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X-ray diffraction experiments at KEK-PF:BL-18C & SPring-8 BL10XU

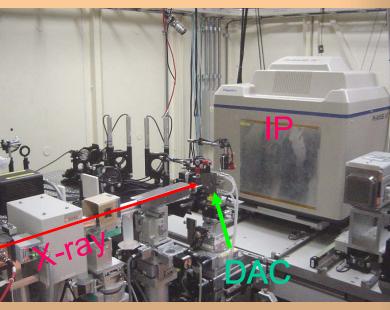
2. Analytical Transmission Electron Microscope equipped with EDS system (Determined by ATEM)

with calibrated K-factors (Cliff and Lorimer, 1975; Fujino et al., 1988) JEOL, JEM-3000F[IMR, Tohoku Univ.]

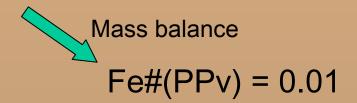
Powder X-ray diffraction



SPring-8, BL10XU



$$Fe\#(Fp) = 0.23$$
 (Fe# = FeO/(MgO+FeO) in the mole fraction)



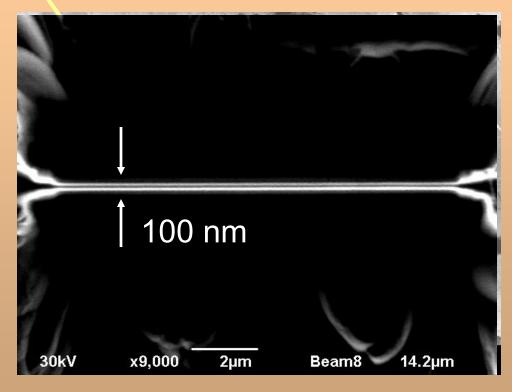


Sample preparation for TEM study by Focused Ion Beam (FIB) system

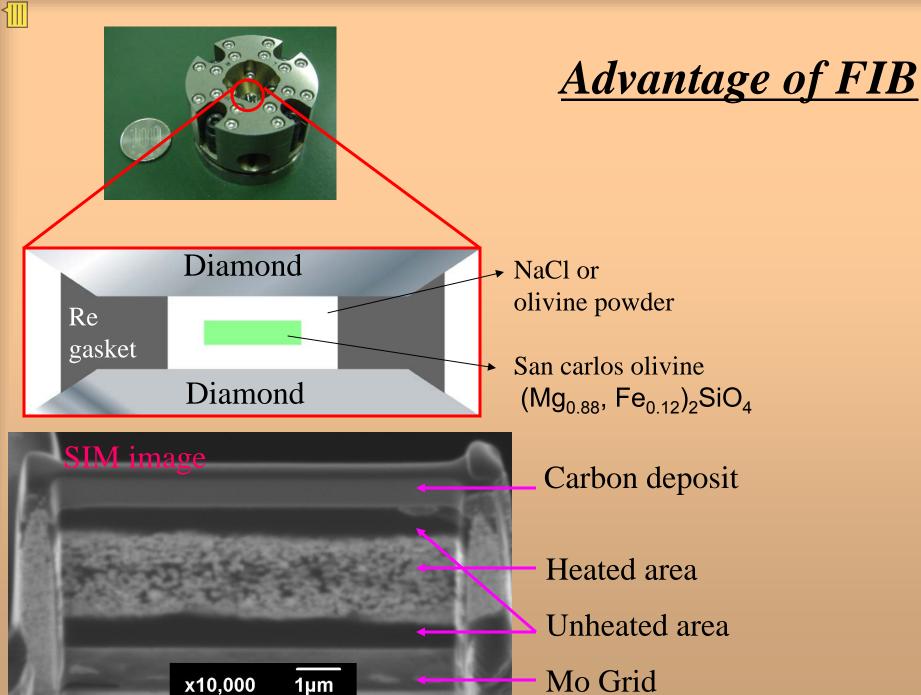


FIB system (Tohoku univ.)

Re gasket

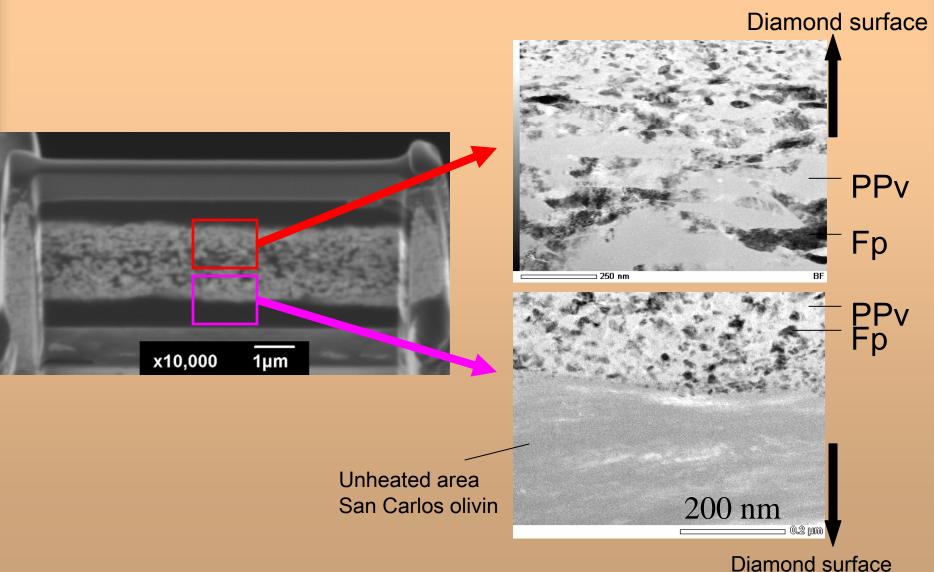


SIM image of recovered sample



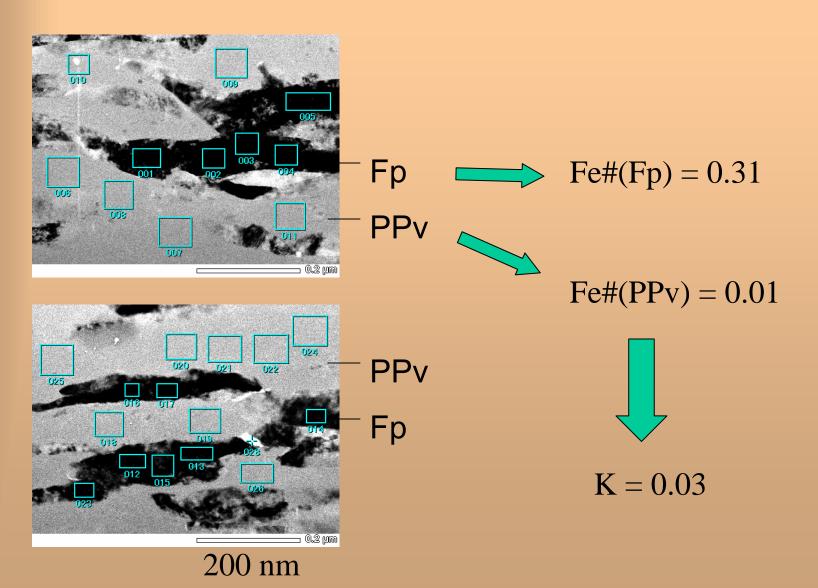
STEM image of the recovered sample

STEM image of the recovered sample from 140GPa, 2000K (90min)

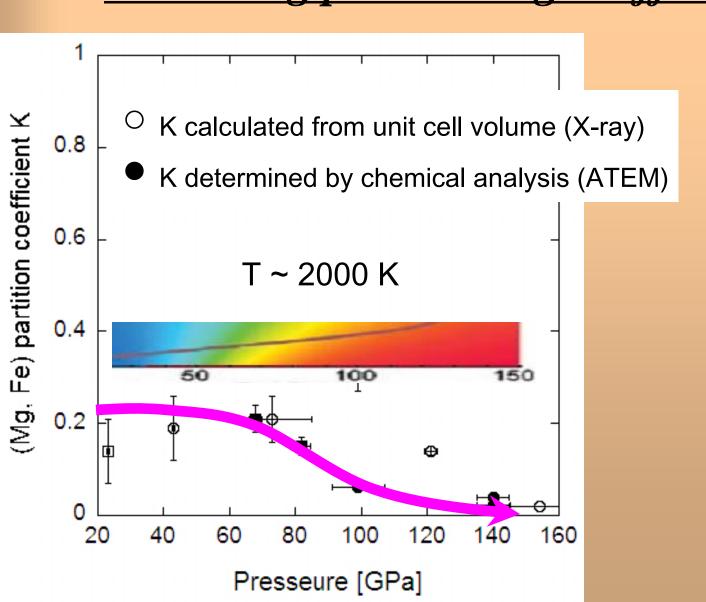


Chemical analysis by ATEM

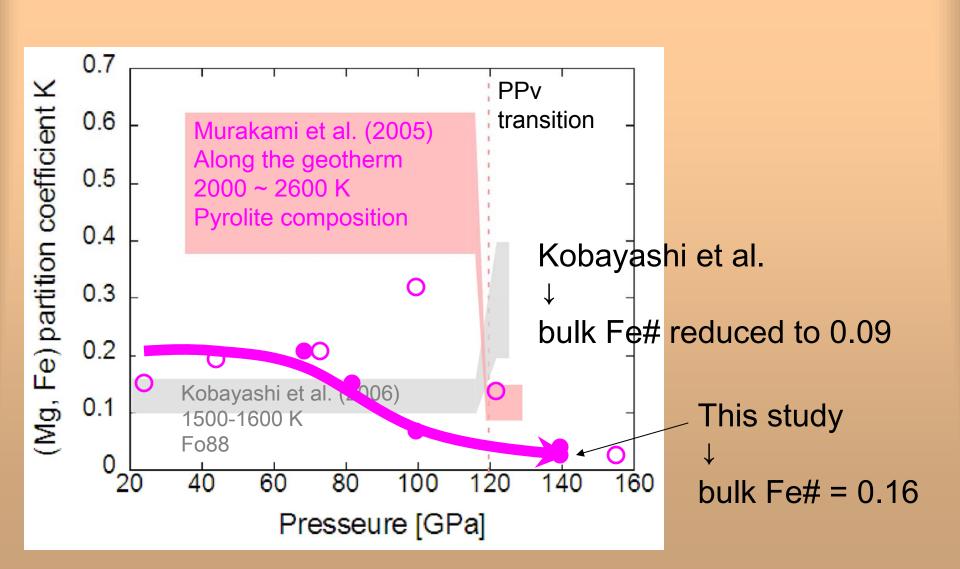
STEM image of the recovered sample from 140GPa, 2000K(90min)



Pressure dependency of the Fe-Mg partitioning coefficient (K)



Partition coefficient (K) of Fe-Mg between Pv/PPv and Fp



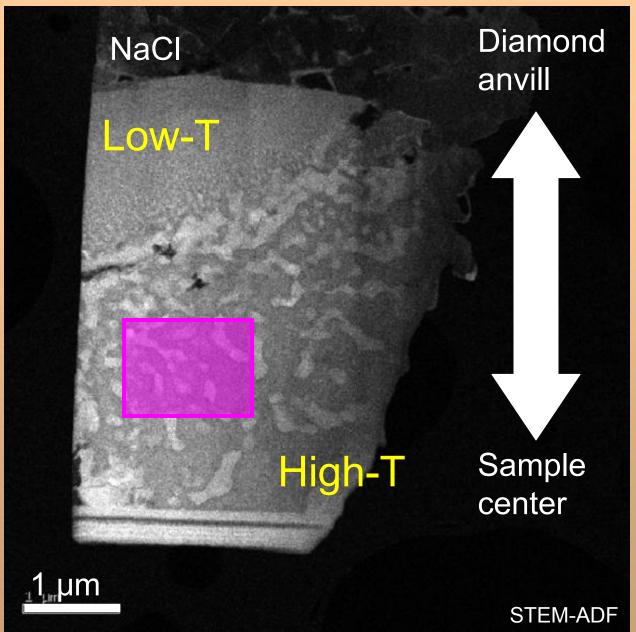
Summary

- ✓ The ATEM analysis shows that post-perovskite phase exhibits with very small iron content of Fe# = 0.01 in the recovered sample from 140 GPa and 2000 K.
- ✓ Iron prefers ferropericlase strongly rather than postperovskite phase, which is consistent with the high-spin/lowspin transition arguments of ferropericlase [Badro et al., 2003].
- ✓ The difference in the partition coefficients between the present result and that by Kobayashi et al. (2005) may be caused by the compositional dependency of the partitioning behavior between post-perovskite and ferropericlase.

title



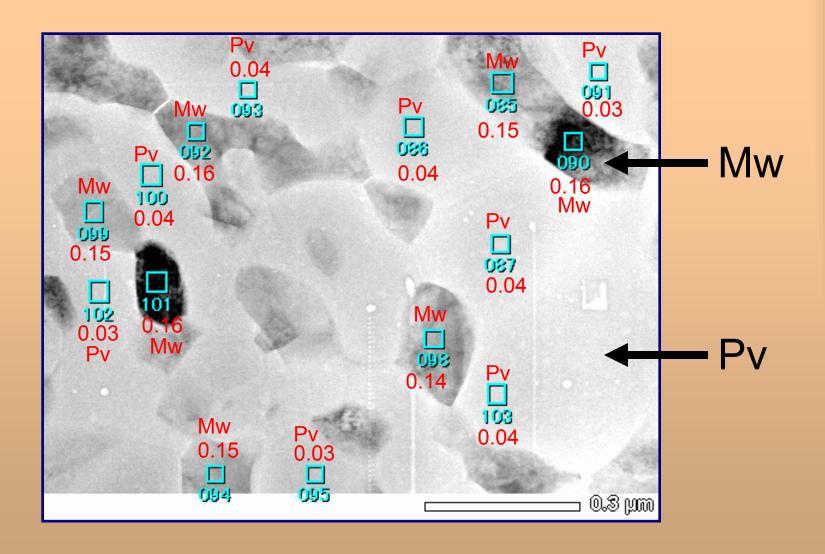
TEM image of the recovered sample



68GPa, 2000K 60min



High magnification image of the sample



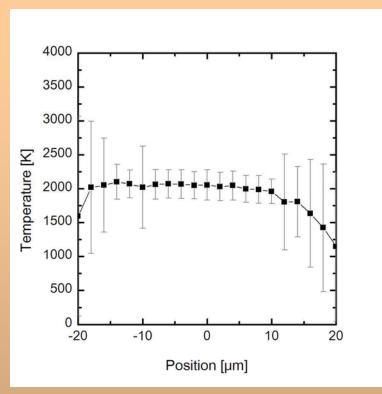


Fig. 2-3 Temperature distribution (Run SC34)

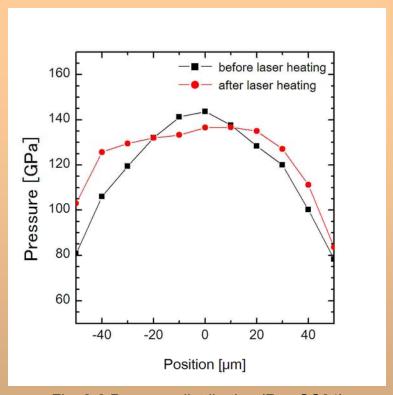


Fig. 2-2 Pressure distribution (Run SC34)

Table 2-1 The composition of San Carlos olivine

San Carlos Olivine								
n=5	wt.%	1σ	at.%	1σ	element cation ratio		1σ	
MgO	47.37	0.10	58.24	0.12	Mg	1.748	0.004	
Al2O3	n.d.	-	n.d	-	Al	-	-	
SiO2	40.37	0.10	33.29	0.08	Si	0.999	0.002	
CaO	0.15	0.03	0.13	0.03	Ca	0.004	0.001	
MnO	0.13	0.07	0.09	0.05	Mn	0.003	0.001	
FeO	11.55	0.09	7.97	0.06	Fe	0.239	0.002	
NiO	0.43	0.04	0.28	0.03	Ni	0.009	0.001	
total	100.00		100.00		Ο	4		
Mg#			0.88		Cation sum.	3.001	0.005	

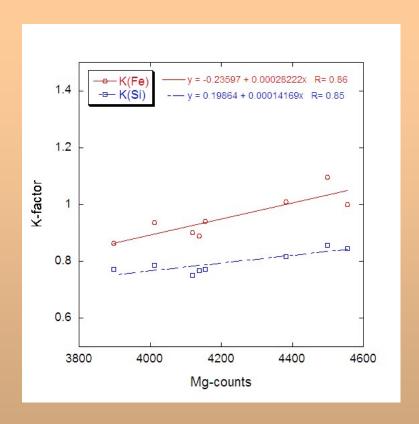
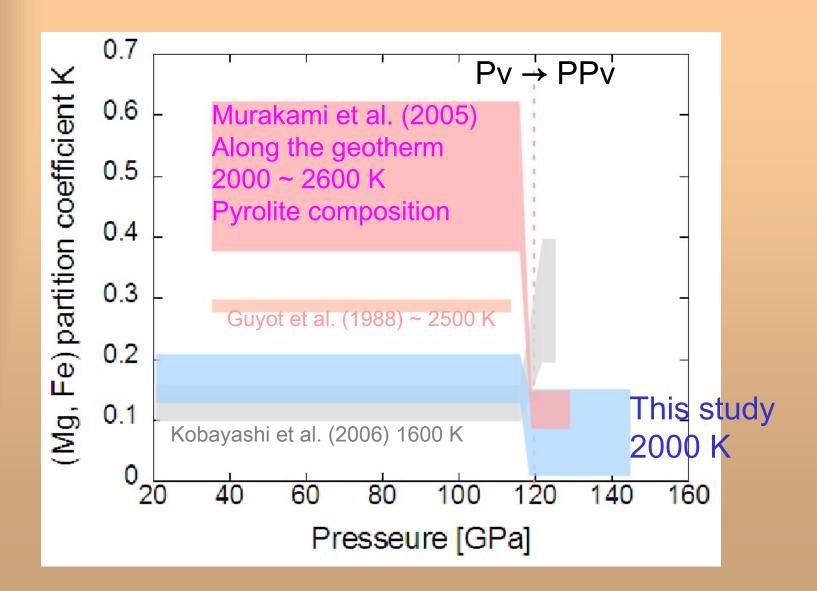


Fig. 3-4 The calibration curve of the k-factor of k_{FeMg} and k_{SiMg}

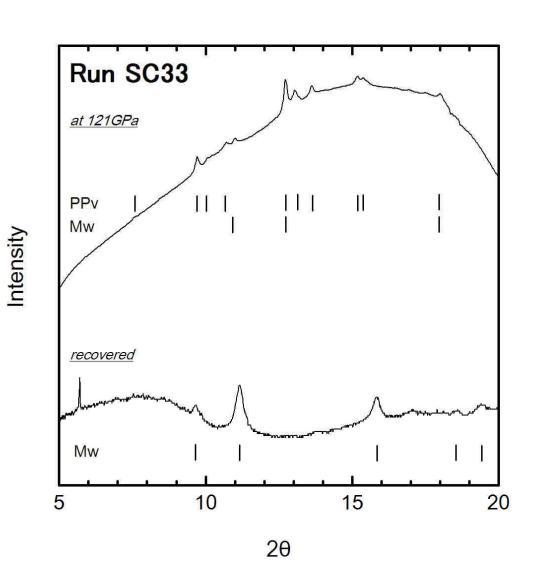
	Post-perovs	kite	Ferropericla	ase	Unheated area		
	n=14		n=12		n=8		
wt.%							
MgO	44.34	2.55	55.12	1.38	47.70	0.52	
Al_2O_3	n.d.	n.d.	n.d.	n.d.	n.d.	n.d.	
SiO ₂	54.50 4.25		(sub)	(sub)	40.66	0.63	
FeO	1.16	0.69	44.88	2.46	11.64	0.36	
total	100.00 5.00		100.00	100.00 2.82		0.89	
at.%							
MgO	54.37	3.12	68.64	1.72	58.52	0.64	
Al_2O_3	n.d.	n.d.	n.d.	n.d.	n.d.	n.d.	
SiO ₂	44.83	3.50	(sub)	(sub)	33.47	0.52	
FeO	0.80	0.47	31.36	1.72	8.01	0.25	
total	100.00	4.71	100.00	2.43	100.00	0.00	
Fe#	0.01	0.01	0.31	0.02	0.12	0.01	

Table 3-3 The composition of post-perovskite and ferropericlase (Run SC34)





Powder X-ray diffraction pattern



121(2)GPa 2000(200)K 60min After laser heating

recovered

Run		P [GPa]	Pv or PPv Fe	# err	Mw Fe#	err	Bulk Fe#	err	K	err
SC44	X- ray	43	0.03	0.01	0.12	0.02	0.07	0.02	0.19	0.07
SC40	X- ray	73	0.04	0.00	0.16	0.03	0.10	0.03	0.21	0.05
SC45	X- ray	99	0.08	0.01	0.22	0.01	0.15	0.01	0.32	0.05
SC33	X- ray	121	0.03		0.21	0.02	0.12	0.02	0.14	0.01
SC34	ATEM	140	0.01	0.01	0.31	0.03	0.16	0.03	0.02	0.02
SC41	X- rav	154	0.01		0.23	0.01	0.12	0.02	0.02	0.00